

# 12.005 Lecture Notes 27

## Flow in Porous Media

Problem of great economic importance (also scientific)

- hydrology (ground water migration, toxic waste)
- oil migration
- soil stability, fault mechanics (pore pressure)
- melt migration in mantle
- geysers and hot springs

Porous medium  $\Rightarrow$  voids  $\Rightarrow$  porosity  $\phi$

$\phi \equiv$  volume fraction of voids

For example,

Sand:  $\phi \sim 40\%$

Pumice:  $\phi \sim 70\%$

Oil shales:  $\phi \sim 10\text{--}20\%$

If pore connected  $\Rightarrow$  permeable

Pressure gradient  $\Rightarrow$  flow

Darcy's law  $\Rightarrow \underline{v} = -\frac{k}{\eta} \nabla p$

$v \equiv$  volumetric flow rate     $k \equiv$  permeability

We can use Poiseuille flow for simple geometries. For example, cubical matrix, circular tubes or pipes.

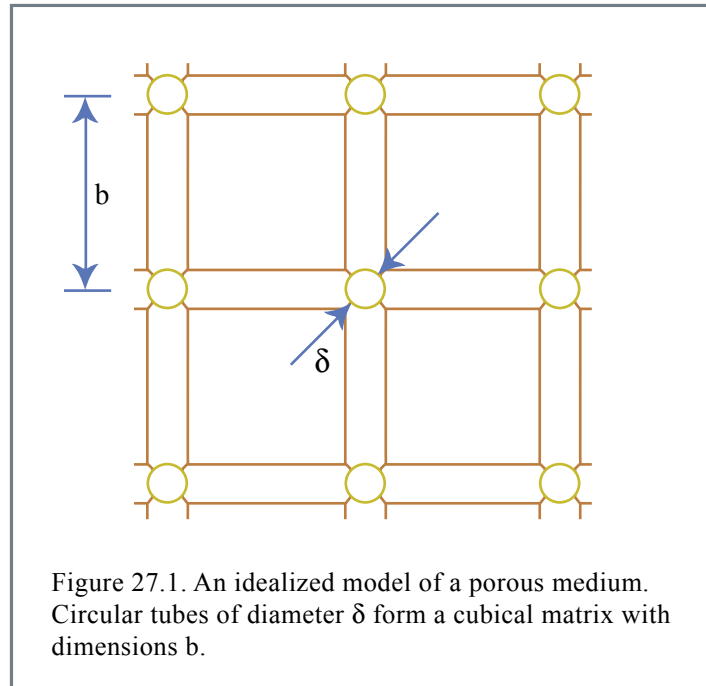


Figure by MIT OCW.

$$\phi = \frac{12 \cdot \frac{1}{4} \cdot \pi \cdot \left(\frac{\delta}{2}\right)^2 \cdot b}{b^3} = \frac{3\pi \delta^2}{4 b^2}$$

Consider  $\frac{dp}{dx}$  (one direction only)

In each pipe (along  $x$ ),  $\bar{u} = -\frac{\delta^2}{32\eta} \frac{dp}{dx}$  [Poiseuille flow]

$$\text{Darcy velocity: } v = \frac{4 \cdot \frac{1}{4} \cdot \bar{u} \cdot \pi \cdot \left(\frac{\delta}{2}\right)^2}{b^2} = \frac{\pi \delta^2}{4b^2} \bar{u} = \frac{\phi}{3} \bar{u}$$

$$v = -\frac{b^2 \phi^2}{72\pi\eta} \frac{dp}{dx}$$

$$\Rightarrow k = \frac{1}{72\pi} b^2 \phi^2$$

Large  $b \Rightarrow$  large  $v$ ?  $b^2 = \frac{3\pi}{4} \frac{\delta^2}{\phi}$

Large  $\phi \Rightarrow$  large  $v$ ?  $k = \frac{\pi}{128} \frac{\delta^4}{b^2}$

Compare to cubes separated along faces (channel flow)

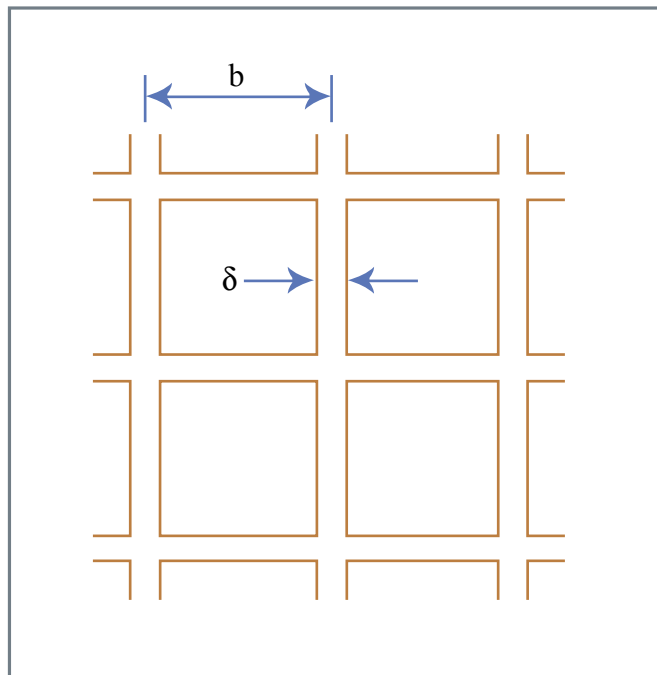


Figure 27.2  
Figure by MIT OCW.

$$\phi = \frac{6 \cdot \frac{1}{2} \cdot \delta b^2}{b^3} = 3 \frac{\delta}{b}$$

Again,  $\frac{dp}{dx}$  directed along one edge

$$u = \frac{1}{2\eta} \frac{dp}{dx} (Z^2 - (\delta/2)^2)$$

$$\bar{u} = \frac{1}{2\eta\delta} \frac{dp}{dx} \left( \frac{Z^3}{3} - \frac{\delta^2 Z}{2} \right) \Big|_{-\delta/2}^{\delta/2} = -\frac{5\delta^2}{24\eta} \frac{dp}{dx}$$

Darcy velocity:  $v = 2 \frac{b\delta}{b^2} \bar{u} = -\frac{5}{12} \frac{\delta^3}{b\eta} \frac{dp}{dx} = -\frac{5}{324} \frac{b^2 \phi^3}{\eta} \frac{dp}{dx}$

$$k = \frac{5b^2 \phi^3}{324}$$

$k$  is different depending on  $\phi$ .

$$k = \frac{135}{324} \frac{\delta^3}{b}$$

Clearly, porosity distribution is important.

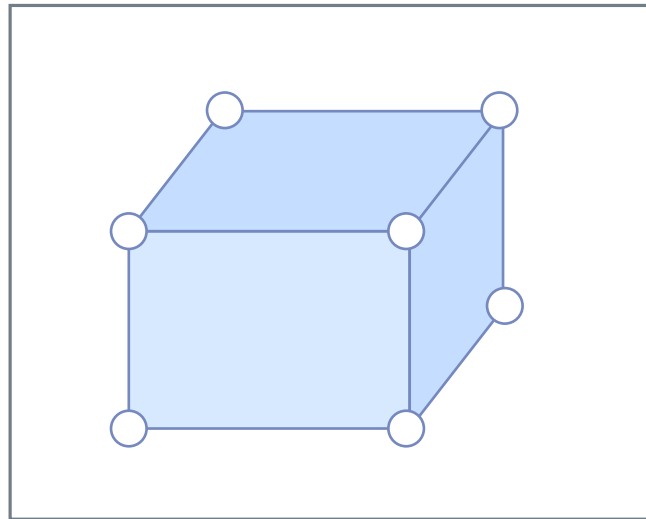


Figure 27.3  
Figure by MIT OCW.

Also -- more easily measured than figured out theoretically -- more complicated geometries → numerical simulation.

Consider “Lawn Sprinkler” example – flow in unconfined aquifer.

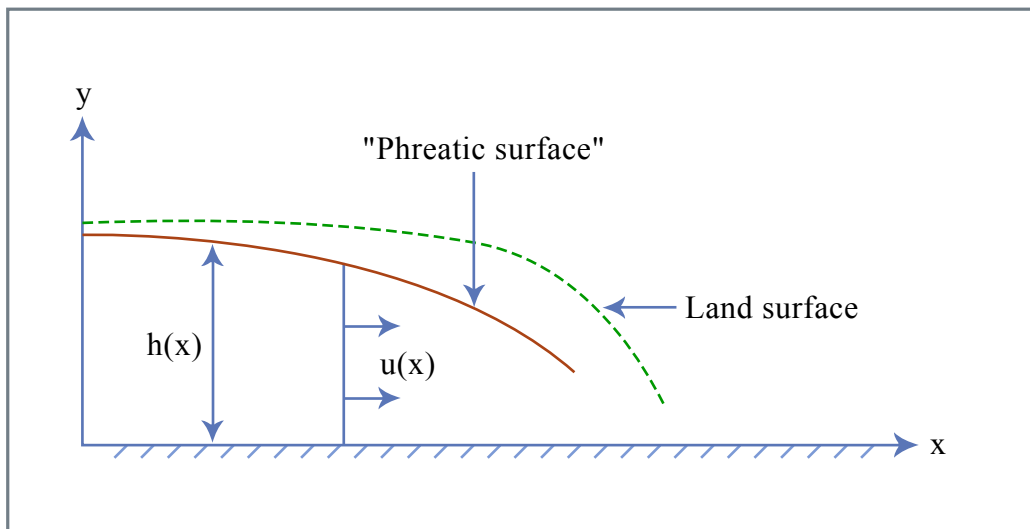


Figure 27.4  
Figure by MIT OCW.

$h \equiv$  “hydraulic head”

$u \rightarrow$  Darcy velocity

Dupuit approximation:  $\frac{dp}{dx} = \rho g \frac{\partial h}{\partial x}$

For  $\frac{\partial h}{\partial x} \ll 1$  flow is one-dimensional.

Darcy’s law:  $u = -\frac{k\rho g}{\eta} \frac{\partial h}{\partial x}$

Conservation of mass: Assume no input

Flux  $Q = u(x)h(x) = -\frac{k\rho g}{\eta} h \frac{dh}{dx} = \text{const.}$

$\Rightarrow$  phreatic surface is a parabola

For  $h = h_0$  at  $x = 0$

$$h = \left( h_0^2 - \frac{2Q\eta x}{k\rho g} \right)^{1/2}$$

Suppose we have a porous dam of width  $w$ . The relation between  $Q$ ,  $h_0$  and  $h_1$  is:

$$Q = \frac{k\rho g}{2\eta w} (h_0^2 - h_1^2)$$

or

$$Q = \frac{k\rho g}{2\eta w} [(h_0 - h_1)(h_0 + h_1)]$$

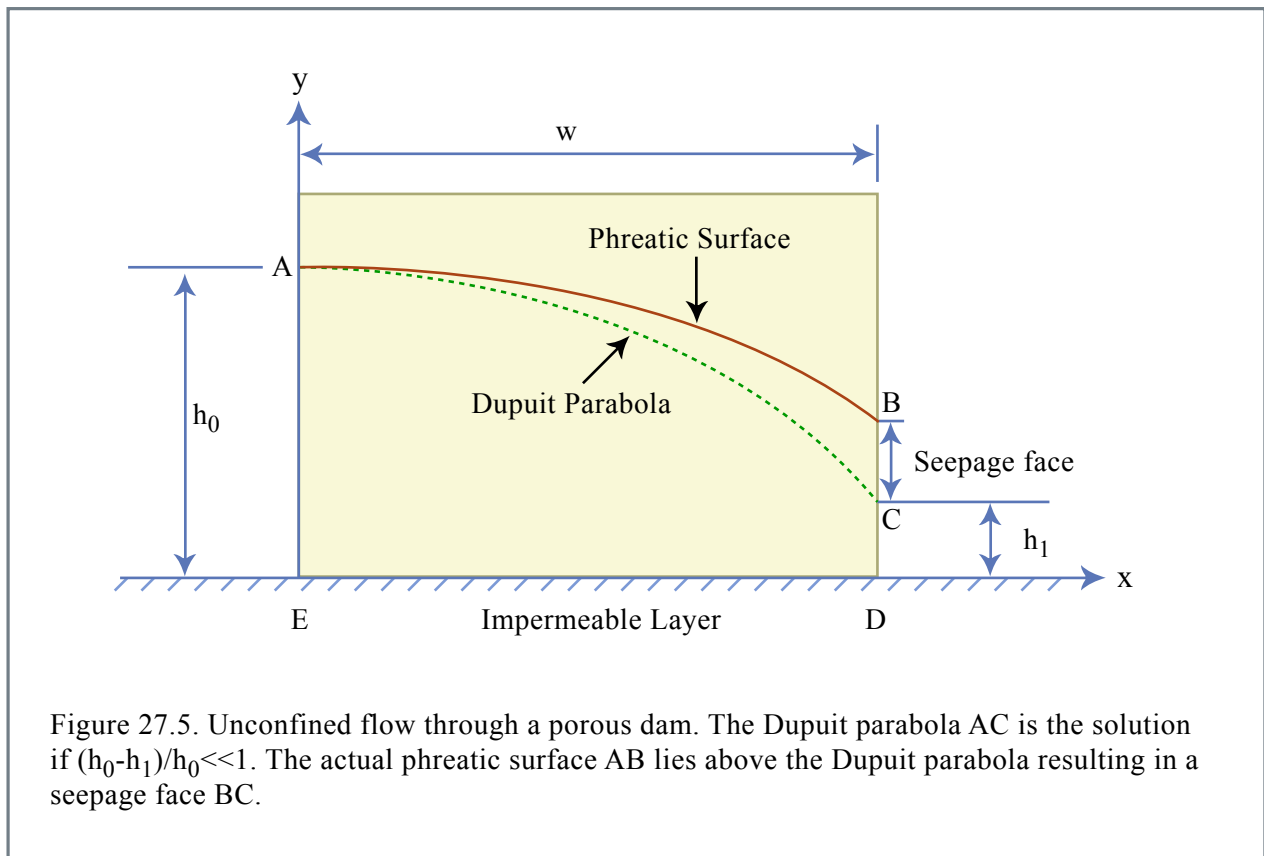


Figure 27.5. Unconfined flow through a porous dam. The Dupuit parabola AC is the solution if  $(h_0 - h_1)/h_0 \ll 1$ . The actual phreatic surface AB lies above the Dupuit parabola resulting in a seepage face BC.

Figure by MIT OCW.