

IV-3 COEFFICIENT OF PERMEABILITY (Hydraulic Conductivity)

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13,782 500 SHEETS FULLER 5 SQUARE
42,781 50 SHEETS EYE-EASE 5 SQUARE
42,382 100 SHEETS EYE-EASE 5 SQUARE
42,383 200 SHEETS EYE-EASE 5 SQUARE
42,384 100 SHEETS EYE-EASE 2 SQUARE
42,385 200 SHEETS EYE-EASE 2 SQUARE
42,386 200 RECYCLED WHITE 2 SQUARE
Made in U.S.A.



COEF. OF PERMEABILITY (Hydraulic Conductivity)

1. THEORETICAL RELATIONSHIPS

1.1 Kozeny - Carman Equation

1) Poiseuille eqn. for laminar flow through a capillary tube



v' = actual velocity

i' = actual gradient

$$v' = \left(\frac{\gamma}{\mu}\right) C_s R_H^2 i'$$

↑ Permeant

• γ = unit weight of fluid (g/cc)

• μ = viscosity of fluid

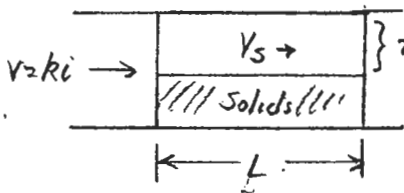
$\approx 10^{-5}$ g-sec/cm² for H₂O at 20°C

• C_s = shape factor of tubes ($\approx 1/3 - 1/2$)

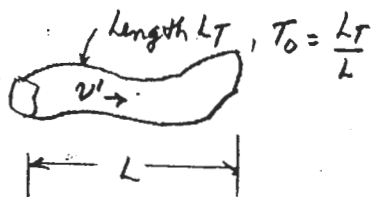
$$R_H = \text{hydraulic radius} = \frac{\text{area of flow}}{\text{wetted perimeter}} = \frac{\text{pore volume}}{\text{Surface area}}$$

2) Convert from capillary tube to soil à la Part IV-2 for

flow per unit area, $Q/A = v = ki$; k = coefficient of permeability
= hydraulic conductivity



$$v = n v_s = \frac{e}{1+e} v_s$$



$$v_s = \frac{v'}{T_0} ; i' = \frac{i}{T_0}$$

$$R_H = \frac{e v_s}{\text{Surface Area}} = \frac{e}{\frac{\text{Surface Area}}{v_s} = \text{SSA} \cdot W_s} = \frac{e}{\text{SSA} \cdot G_s} \quad (\text{for } \gamma_w = 1)$$

3) Resultant equation

$$v = ki = \frac{e}{1+e} \left[v_s = \frac{v'}{T_0} = \left(\frac{\gamma}{\mu}\right) \frac{C_s}{T_0} \left(\frac{e}{\text{SSA} \cdot G_s}\right)^2 \frac{i}{T_0} \right]$$

$$\therefore k \text{ (cm/sec)} = \left(\frac{\gamma}{\mu}\right) \left[\frac{C_s}{T_0^2} \left(\frac{e}{1+e}\right) \left(\frac{e}{\text{SSA} \cdot G_s}\right)^2 \right]$$

Permeant
($\frac{1}{\text{cm} \cdot \text{sec}}$)
Physical Permeability, K
(cm²)

1.2 Effect of Permeant

- 1) If no change in fabric, then k affected only by γ/μ
- 2) For flow of water at 20°C, $k(\text{cm/sec}) \approx 10^5 (K, \text{cm}^2)$
- 3) Effect of temperature

$$k_T = k_{20^\circ C} \frac{\mu_{20^\circ}}{\mu_T}$$

	Reference (Lab)		
	↓		
T(°C)	10	20	30
μ (milli poise)	13	10	8
	x 0.77		x 1.25

• $\approx 25\%$ change / 10°C

• In situ T = 10°C NE

1 poise = $\frac{\text{dyne} \cdot \text{sec}}{\text{cm}^2}$; poise $\times 10^{-1} = \text{Pa} \cdot \text{sec}$

2. HYDRAULIC CONDUCTIVITY OF SOILS (flow of Water)

2.1 General Magnitudes [Also see Fig. 19.5 of L/W ('69)]

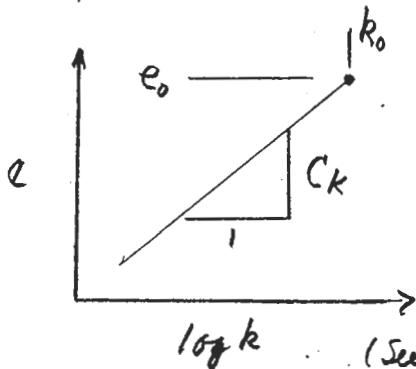
Soil Type	$k(\text{cm/sec})$
Fine sand	10^{-2} $k \approx (D_{10\text{mm}})^2$
Non-plastic silt	10^{-5}
CL-CH clays	$10^{-8 \pm 1}$

2.2 Effect of Void Ratio

- 1) Granular soils

$$k \propto \frac{e^3}{(1+e)} \cdot S^3 \text{ with approx. relationship for } S < 100\%$$

- 2) Saturated, natural cohesive soils



$$C_k = \frac{de}{d \log k} \approx \left(\frac{1}{3} - \frac{1}{2}\right) e_0$$

(see Sheet A, Fig. 17)

$$\therefore k = k_0 (10)^{\left(\frac{e - e_0}{C_k}\right)}$$

(See Sheet A, Fig. 11 for example)



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2.3 Effect of Fabric with Sedimentary Clays



Flow mainly controlled by voids between floccs

- Flocc size = function particle size
 { shape + environment Part II-2

$\left(\frac{e}{SSA \cdot G_s}\right)^2$ term should be for floccs, not individual particles

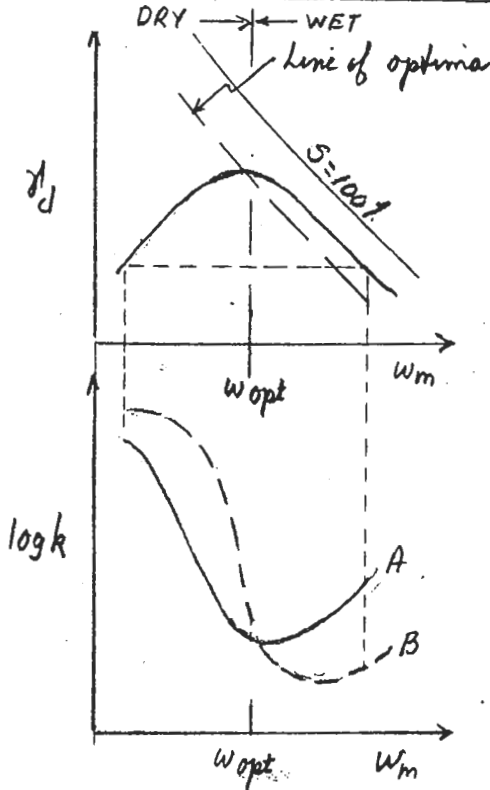
- Clays do not require an "critical" gradient for flow
- For marine illitic clays, $\gamma_k = k_h/k_v = 1.0-1.5$ à la Fig 12, sheet A
 Also $k_v = 10^{-8} \text{ cm/sec}$

2.4 Compacted Cohesive Soils

2.4.1 Overview of Compaction

See Sheet B

2.4.2 Variation in k with w_m & γ_d (Dynamic or kneading compaction)



a) Importance of fabric

- Dry of optimum \rightarrow dense aggregates with large voids \rightarrow high k FLOCCULATED FABRIC
- Wet of optimum \rightarrow more uniform distribution of particles w/ small voids \rightarrow low k DISPERSED FABRIC
- \therefore at same γ_d , $k_{wet} \ll k_{dry}$

b) Minimum k for given Comp. Effort

- Curve A $\rightarrow k_{min}$ at $w_m = w_{opt}$, i.e. at max. γ_d (min. e)
- Curve B $\rightarrow k_{min}$ at $w_m > w_{opt}$, i.e., Usual Δ fabric more important than dec. γ_d Case

* \therefore In field, always use $w_m \geq w_{opt}$ to get low k for clay liner *

c) See Sheet C for actual data

3. LABORATORY MEASUREMENT TECHNIQUES

3.1 Natural Cohesive Soils

Note • Tests on specimens trimmed from undisturbed tube samples

- Measure k as $f(\sigma'_c, \sigma'_{vc}) \rightarrow e$ vs $\log k$ data

1) Triaxial cell

- Requires top & bottom porous stones w/ drainage lines
- Usually run constant head test (with u back pressure) with burettes to measure both inflow & outflow (check for leaks)

2) Modified oedometer with falling head test

See Tavenas et. al. (1983) CGJ, 20(4), Fig. 8 - Oedometer cell adapted for falling head permeability tests.

3) CRS consolidation

$$\dot{\epsilon} = d\epsilon/dt$$

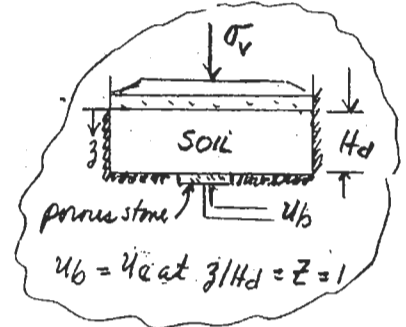
$$\epsilon = \epsilon_v = \Delta H/H_0$$

- Apply constant $\dot{\epsilon}$ with measurements of σ_v , u_b and ϵ_v (want $u_b/\sigma_v \approx 20 \pm 10\%$)

• For constant $m_v = d\epsilon_v/d\sigma'_v$

(a) Ave. $\sigma'_{vc} = \sigma_v - \frac{2}{3} u_b$

(b) $k_v = \frac{\dot{\epsilon} H_d^2 \gamma_w}{2 u_b}$



$u_b = u_c$ at $z/H_d = z = 1$

NOTE: Equations also available for constant $CR = d\epsilon_v/d \log \sigma'_{vc}$

Part IV-2 Notes

(c) Coef. of consolidation, $c_v = \frac{k_v}{m_v \gamma_w} = \frac{H_d^2}{2 u_b} \left(\frac{d\sigma_v}{d\epsilon} \right)$
 $t = \dot{\epsilon} dt / d\sigma_v$

3.2 Compacted Cohesive Soils

Comments

- a) Compaction mold, ^{used} only for design (since compact soil in mold)
- b) Triaxial cell most common for testing tube samples of compacted soil

13 762
42 381
42 386
42 392
42 399
42 389
500 SHEETS, FILLER, 5 SQUARE
50 SHEETS, LVC CASE, 5 SQUARE
100 SHEETS, LVC CASE, 5 SQUARE
100 SHEETS, LVC CASE, 5 SQUARE
100 RECYCLED WHITE, 5 SQUARE
100 RECYCLED WHITE, 5 SQUARE
MADE IN U.S.A.

